

## ENTHALPIES OF FORMATION OF MINERALS OF SERPENTINE GROUP

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The minerals of magnesium serpentine group of composition  $Mg_3Si_2O_5(OH)_4$ : natural scale antigorite (Saronovskoe deposit, N.-W. Ural), chrysotile-asbestos from serpentinized ultrabasites (E. Sayani) and synthesized under hydrothermal conditions lizardite were studied. The diagnostic investigations were performed by electron microprobe and thermal analyses, X-ray diffraction and IR-spectroscopy. For thermochemical study of serpentines the high-temperature heat-flux Tian-Calvet microcalorimeter was used. The enthalpies of formation were determined by “transposed temperature drop solution calorimetry” method. A samples of 1.5-10 ( $\pm 2 \times 10^{-3}$ ) mg were thermostated at room temperature and were thrown down into the calorimeter with a melt  $2PbO \cdot B_2O_3$  as a solvent at  $T=973$  K. The heat measured contained two contributions: the heat content of the samples and the enthalpy of solution at  $T=973$  K -  $[H^0(973 \text{ K})-H^0(298.15 \text{ K})+\Delta_{sol}H^0(973 \text{ K})]$ . The calorimeter was calibrated in the same manner by using a platinum reference sample with known enthalpy increments [1]. Using experimental and needed thermochemical data for constituent oxides the standard enthalpies of formation of minerals studied were calculated (table). The values of  $\Delta_f H^0_{el}(298.15 \text{ K})$  obtained are in a good agreement within experimental errors. Differences between magnesium serpentines are associated with some disproportions between octahedral and tetrahedral nets in structure, which lead to formation of different structural lays - planar (lizardite), cylindrical (chrysotile) и curved modulated (antigorite). Our results show minor energetic distinctions between these three structural types.

**Table**

Thermochemical data for minerals studied and constituent oxides (kJ/mol)

| Substance            | $H^0(973 \text{ K})-H^0(298.15 \text{ K})+\Delta_{sol}H^0(973 \text{ K})$ | $-\Delta_f H^0_{ox}(298.15 \text{ K})$ | $-\Delta_f H^0_{el}(298.15 \text{ K})$ |
|----------------------|---|--|--|
| MgO (s)              | $36.38 \pm 0.59^a$  |  | $601.6 \pm 0.3$ [1]                    |
| SiO <sub>2</sub> (s) | $39.43 \pm 0.21^b$  |  | $910.7 \pm 1.0$ [1]                    |
| H <sub>2</sub> O (l) | $40.9 \pm 2.5^c$  |  | $285.8 \pm 0.1$ [1]                    |
| Antigorite           | $445.1 \pm 10.6$ (8) <sup>d</sup>   | $170 \pm 12$                           | $4368 \pm 12$                          |
| Chrysotile           | $425.7 \pm 16.5$ (6) <sup>d</sup>   | $156 \pm 17$                           | $4353 \pm 17$                          |
| Lizardite            | $425.5 \pm 10.2$ (4) <sup>d</sup>   | $156 \pm 12$                           | $4353 \pm 12$                          |

<sup>a, b</sup> – calculated using reference data on  $[H^0(T)-H^0(298.15 \text{ K})]$  [1] and experimental data on  $\Delta_{sol}H^0(973 \text{ K})$ : <sup>a</sup> – [2], <sup>b</sup> – [3]; <sup>c</sup> – estimated in [4]; <sup>d</sup> - errors are expressed by interval for P=95%, in brackets the number of determinations is given

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